

# Mean Tropospheric Residence Time of Cosmic-Ray-Produced Beryllium 7 at North Temperate Latitudes

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By means of published data on the concentration of  ${}^7\text{Be}$  in the troposphere between  $30^\circ$  and  $75^\circ\text{N}$ , together with concentrations measured over the Netherlands, the tropospheric content of  ${}^7\text{Be}$  has been estimated. Comparison of this value with published rates of deposition of  ${}^7\text{Be}$  in the same latitude band shows that the mean tropospheric residence time of this nuclide is probably not less than 22 days and is more likely of the order of 35 days. This is in agreement with previous estimates and not in agreement with the value of 1 week or less proposed recently by Martell and Moore.

## INTRODUCTION

In a discussion of published estimates of the tropospheric residence time of aerosols carrying natural and artificial radioactive nuclides, *Martell and Moore* [1974] concluded that the mean tropospheric residence time of these aerosols is less than 1 week. Longer residence times, including those proposed for cosmic ray spallation products, were assumed to be in error.

The mean tropospheric residence time  $\tau$  of cosmic-ray-produced nuclides can be deduced from measurements of the average tropospheric content  $N$  (atoms  $\text{m}^{-2}$ ) and the mean deposition rate  $D$  (atoms  $\text{m}^{-2} \text{s}^{-1}$ ) on the earth's surface ( $\tau D = N$ ). It may also be calculated from the known tropospheric production rate  $P$  and the measured deposition rate ( $P = D(1 + \lambda\tau)$ , where  $\lambda$  is the disintegration constant of the nuclide) or from the ratio of the tropospheric concentrations of two nuclides with different half-lives, the residence times for the two being assumed to be identical.

At higher latitudes the last two procedures are not applicable because exchange of air between stratosphere and troposphere, which may influence the result appreciably, is not taken into account. Influx of stratospheric  ${}^7\text{Be}$  down to an altitude of 3 km has been shown to occur by *R. Reiter et al.* [1971]. Assuming a mean tropospheric residence time of 30 days, *Bleichrodt and van Abkoude* [1963b] estimated that some 30% of the deposition of  ${}^7\text{Be}$  at  $50^\circ\text{N}$  originates from the stratosphere; for a residence time of 7 days this percentage would be about 40. *Walton and Fried* [1962] showed that the highest value for the radioactivity ratio of  ${}^7\text{Be}$  and  ${}^{32}\text{P}$  observed in 14 rains at Westwood ( $41^\circ\text{N}$ ,  $74^\circ\text{W}$ ) might, for instance, result from exclusive irradiation of an air parcel within the troposphere for 68 days or from the presence of 12% of stratospheric air in a tropospheric air sample which was cleansed by rain 10 days before. A small percentage of stratospheric air might thus have a great influence on the residence time as estimated from nuclide ratios.

The first method of calculation mentioned above remains applicable if the stratosphere contributes significantly to the tropospheric inventory, because it is based on the definition of residence time, i.e.,  $\tau D = N$ . It therefore seemed worthwhile to reestimate the mean tropospheric residence time of a cosmogenic nuclide by using this method and to compare the result with the value suggested by *Martell and Moore* [1974].

## TROPOSPHERIC CONTENT OF ${}^7\text{Be}$

The amount of  ${}^7\text{Be}$  in a column of air extending from the earth's surface to the tropopause can be computed when the

concentration of  ${}^7\text{Be}$  in air is known as a function of altitude. To construct a vertical concentration profile, as many data as possible have been collected (including those in Table 1) and plotted in Figure 1. Data for the same altitude and groups of data covering small altitude ranges were averaged. The average concentration of  ${}^7\text{Be}$  in air at sea level was taken from Figure 2.

The question may be asked whether the average  ${}^7\text{Be}$  concentration measured over land (Figure 2) equals that over the sea. On the average, the concentration of  ${}^{90}\text{Sr}$  from nuclear tests was found not to differ significantly in surface air over land and ocean between  $35^\circ\text{N}$  and  $55^\circ\text{N}$  during the years 1966–1967 [*Freudenthal*, 1969]. Since the  ${}^{90}\text{Sr}$  originated mainly from the stratosphere, from which injection into the troposphere depends on latitude, whereas a large part of the  ${}^7\text{Be}$  in surface air has been produced in the troposphere with little horizontal variation in production rate [*Lal and Suess*, 1968], any difference in the concentrations over land and sea is probably smaller for  ${}^7\text{Be}$  than for  ${}^{90}\text{Sr}$ .

The average values in Figure 1 show so much scatter that it may be doubted whether drawing an average concentration profile will lead to a realistic estimate of the tropospheric content of  ${}^7\text{Be}$ . To determine a minimum value for the tropospheric content and thus a minimum for the residence time, profile I in Figure 1 was drawn after rejection of the highest value of each group of data of which the average had been plotted. Curves II and III will certainly yield much closer approximations of the true tropospheric content.

By means of the curves in Figure 1 and data on the 1962 COESA (U.S. Committee on Extension to the Standard Atmosphere) standard atmosphere [*Sissenwine*, 1969] the amount of  ${}^7\text{Be}$  in a column of air in the troposphere has been calculated by plotting the concentration against altitude in kilograms per square meter of pressure and integrating graphically between ground level and various tropopause heights. The results are shown in Table 2.

The number of  ${}^7\text{Be}$  data obtained by aircraft did not allow an estimation of concentration profiles for different seasons. To minimize the influence of seasonal variations in the concentration of  ${}^7\text{Be}$  in surface air and in the deposition of  ${}^7\text{Be}$  (see discussion below), data have been averaged as much as possible for periods of a whole number of years.

## AVERAGE DEPOSITION

Published values of the total deposition of  ${}^7\text{Be}$ , i.e., the sum of dry and wet fallout, per unit of precipitation between  $30^\circ$  and  $70^\circ\text{N}$ , have been collected in Table 3. The average for the

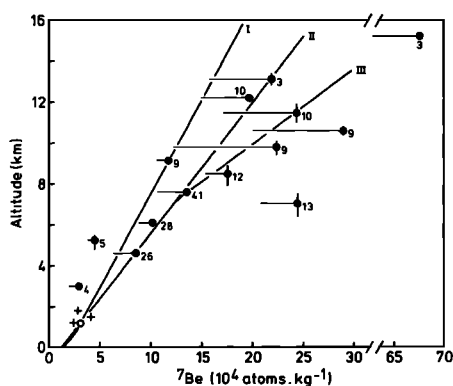


Fig. 1. Concentration of  ${}^7\text{Be}$  in the troposphere. Data points for altitudes above 2 km represent averages of  ${}^7\text{Be}$  concentrations in air sampled by aircraft. The data were taken from Feely *et al.* [1965, 1966], Hardy *et al.* [1961a, b, c, 1964, 1966a, b, 1967, 1968, 1969, 1970, 1972] (only data with a standard error of  $<20\%$ ), Rama and Honda [1961], Bhandari and Rama [1963], Young *et al.* [1970], Winiger *et al.* [1976], and Table 1. All the data pertain to the latitude band  $30^\circ$ – $75^\circ\text{N}$ . To reduce the chance of including results of samples taken partly in the stratosphere, data bearing on altitudes higher than 300 m below the tropopause have been excluded. In case no tropopause level has been reported for samples, only those taken below 8 km have been included. When data for different altitudes were averaged, the range of altitudes is represented by a vertical bar. A horizontal bar connects the average concentration with the average obtained after rejecting the highest value of the corresponding group of data. The numbers indicate the number of data averaged. The open circle represents the average of the data of Peirson [1963] for samples taken at 1.2 km about five times a week during the period October 1959 through August 1976 [Hardy, 1977], are represented by plus signs. The value ( $47^\circ\text{N}$ ,  $112^\circ\text{W}$ ; 1187 m, October 1973 through August 1976), Salt Lake City, Utah ( $41^\circ\text{N}$ ,  $111^\circ\text{W}$ ; 1516 m, January 1972 through December 1974), and Rocky Flats, Colorado ( $40^\circ\text{N}$ ,  $105^\circ\text{W}$ ; average of stations I (1830 m) and II (1738 m) for period September 1972 through August 1976 [Hardy, 1977]), are represented by plus signs. The value for the concentration at sea level corresponds to the average in Figure 2.

10 stations amounts to  $9 \times 10^6$  atoms/l. Probably the best estimate of the mean deposition rate in a latitude band is the product of this value and the average precipitation rate in that band [cf. UNSCEAR, 1964; Malakhov and Pudovkina, 1970].

The mean annual precipitation in various latitude bands has been derived from the recent estimates of Jaeger [1976]. The resulting annual  ${}^7\text{Be}$  deposition in those bands is given in Table 4.

In connection with this treatment of the deposition data it should be noted that several investigators have observed an inverse correlation between specific radioactivity and amount of precipitation for debris from nuclear test explosions. This phenomenon did not result, however, in a higher annual deposition in areas of low rainfall [UNSCEAR, 1962]. If a similar relation holds for  ${}^7\text{Be}$ , the yearly deposition at the comparatively wet locations where  ${}^7\text{Be}$  deposition was actually measured (Table 3) will exceed that in relatively dry regions. The average value of the deposition of  ${}^7\text{Be}$  at the stations in Table 3, which amounts to  $8 \times 10^6$  atoms/ $\text{m}^2$  yr, will therefore overestimate the mean deposition at temperate latitudes. According to the assumption of a linear relation between amount of deposition and amount of precipitation in a latitude band the annual deposition of  ${}^7\text{Be}$  in the band  $30^\circ$ – $80^\circ\text{N}$  is  $6.2 \times 10^6$  atoms/ $\text{m}^2$  yr (Table 4). This estimate may therefore be too low by maximally 30%. For the other latitude bands the maximum error is smaller (see Table 4).

Several investigations of the concentration of radionuclides in the sea suggest that fallout is higher over oceans than over

continents. This problem has been discussed by Volchok [1974]. The amount of  ${}^{90}\text{Sr}$  produced in nuclear tests appears to balance the amounts accountable in the environment if it is assumed that fallout per unit area over the oceans equals the average fallout measured over land. Also, the stratospheric depletion of  ${}^{90}\text{Sr}$  per year was found to equal the annual amount deposited on the earth's surface. No indications were found of much more precipitation on the ocean than on the adjacent land, of an excess of dry  ${}^{90}\text{Sr}$  fallout in the seas, or of an extra oceanic deposition by capture of  ${}^{90}\text{Sr}$ -bearing aerosol particles by sea spray. Further, it appeared that land runoff could not contribute significantly to the apparent excess  ${}^{90}\text{Sr}$  in the oceans. However, if contrary to this negative evidence, deposition on the sea is indeed a factor of about 2 higher than that on land [Volchok, 1974] and if  ${}^7\text{Be}$  shows a similar behavior [cf. Silker, 1972a, b], the average deposition of this nuclide in the various latitude bands (Table 4) will be higher by a

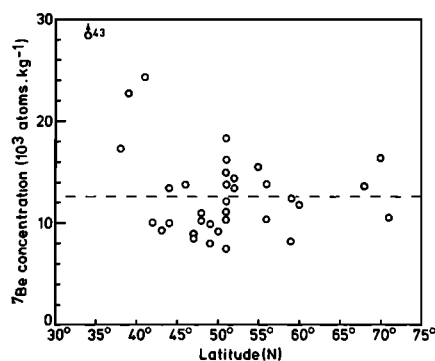


Fig. 2. Yearly average concentrations of  ${}^7\text{Be}$  in surface air at various locations below 200-m altitude. Data for Braunschweig ( $52^\circ\text{N}$ ,  $10^\circ\text{E}$ ; 1970–1974), Lille ( $51^\circ\text{N}$ ,  $3^\circ\text{E}$ ; 1970–1972), Brussels ( $51^\circ\text{N}$ ,  $4^\circ\text{E}$ ; 1973–1974), Mol ( $51^\circ\text{N}$ ,  $5^\circ\text{E}$ ; 1970–1974), Jülich ( $51^\circ\text{N}$ ,  $6^\circ\text{E}$ ; 1974), Cherbourg ( $50^\circ\text{N}$ ,  $2^\circ\text{W}$ ; 1970–1972), Strasbourg ( $49^\circ\text{N}$ ,  $8^\circ\text{E}$ ; 1970–1972), Le Vésinet ( $48^\circ\text{N}$ ,  $2^\circ\text{E}$ ; 1970–1972), Brest ( $48^\circ\text{N}$ ,  $5^\circ\text{E}$ ; 1970–1972), Tours ( $47^\circ\text{N}$ ,  $1^\circ\text{E}$ ; 1970–1972), Bourges ( $47^\circ\text{N}$ ,  $2^\circ\text{E}$ ; 1970–1972), Nîmes ( $44^\circ\text{N}$ ,  $4^\circ\text{E}$ ; 1970–1972), Nice ( $44^\circ\text{N}$ ,  $7^\circ\text{E}$ ; 1970–1972), Biarritz ( $43^\circ\text{N}$ ,  $2^\circ\text{W}$ ; 1970–1972) [Commission of the European Communities, 1972a, b, 1974, 1976]; Kiruna ( $68^\circ\text{N}$ ,  $20^\circ\text{E}$ ; August 1973 through June 1975), Grindsjön ( $59^\circ\text{N}$ ,  $18^\circ\text{E}$ ; August 1972 through June 1975), Ljungbyhed ( $56^\circ\text{N}$ ,  $14^\circ\text{E}$ ; July 1974 through June 1975) [Armsing *et al.*, 1977]; Barrow ( $71^\circ\text{N}$ ,  $157^\circ\text{W}$ ; September 1975 through August 1976), Kap Tobin ( $70^\circ\text{N}$ ,  $22^\circ\text{W}$ ; October 1973 through August 1976), Moosonee ( $51^\circ\text{N}$ ,  $81^\circ\text{W}$ ; September 1970 through August 1976), New York ( $41^\circ\text{N}$ ,  $74^\circ\text{W}$ ; September 1970 through August 1976), Sterling ( $39^\circ\text{N}$ ,  $77^\circ\text{W}$ ; May 1970 through April 1974), Richmond ( $38^\circ\text{N}$ ,  $122^\circ\text{W}$ ; September 1973 through August 1976) [Hardy, 1977]; Leningrad ( $60^\circ\text{N}$ ,  $30^\circ\text{E}$ ; 1965–1968) [Gedeonov *et al.*, 1971]; Stockholm ( $59^\circ\text{N}$ ,  $18^\circ\text{E}$ ; 1960, 1965–1966) [Lindblom, 1962, 1969]; Risö ( $56^\circ\text{N}$ ,  $12^\circ\text{E}$ ; 1966–1967) [Aarkrog and Lippert, 1967, 1968]; Vilnius ( $55^\circ\text{N}$ ,  $25^\circ\text{E}$ ; 48 samples during April 1965 through May 1969) [Luyanas *et al.*, 1970]; Rijswijk ( $52^\circ\text{N}$ ,  $4^\circ\text{E}$ ; November 1960 through October 1961,  $14.4 \times 10^6$  atoms/kg) (J. F. Bleichrodt, unpublished data, 1961); Braunschweig ( $52^\circ\text{N}$ ,  $10^\circ\text{E}$ ; 1964–1968) [Kolb, 1970]; Chilton ( $51^\circ\text{N}$ ,  $1^\circ\text{W}$ ; October 1959 through August 1961, 1964–1965) [Peirson, 1963, also private communication, 1962]; Sutton ( $51^\circ\text{N}$ ,  $0^\circ\text{W}$ ; 30 samples during January 1960 through August 1961) [Parker, 1962]; Heidelberg ( $49^\circ\text{N}$ ,  $9^\circ\text{E}$ ; July 1960 through June 1961) [Schumann and Stoeppler, 1963]; Richland ( $46^\circ\text{N}$ ,  $119^\circ\text{W}$ ; 15 samples during 1967) [Young *et al.*, 1970]; Argonne ( $42^\circ\text{N}$ ,  $88^\circ\text{W}$ ; 1960) [Gustafson *et al.*, 1961], calculated for branching ratio of  ${}^7\text{Be}$  decay of 10.4% [Poenitz and Devolpi, 1973]; and Fullerton ( $34^\circ\text{N}$ ,  $118^\circ\text{W}$ ; September 1973 through July 1975) [Shapiro and Forbes-Resha, 1976]. The dashed line ( $12.7 \times 10^6$  atoms/kg) represents the average when the value of  $43 \times 10^6$  atoms/kg at  $34^\circ\text{N}$  is omitted. (Inclusion of this value will lead to a higher tropospheric residence time.) If the data points are given a weight corresponding with the number of years that air samples were measured, the weighted average amounts to  $13.0 \times 10^6$  atoms/kg.

TABLE 1. Concentration of Beryllium 7 in the Upper Troposphere Over the Netherlands

Date of Sampling	Altitude, km	Tropopause, km	<sup>7</sup> Be, 10 <sup>8</sup> atoms/kg	Maximum Error in Sample Mass, %	Artificial <sup>7</sup> Be, %
Dec. 1, 1960	10.7	12.8	0.88 ± 0.12†	20	
Jan. 25, 1961	8.2	10.4	2.04 ± 0.30	7	
April 20, 1961	8.5	10.2	4.2 ± 0.3	20	
April 27, 1961	8.8	10.2	4.1 ± 0.1	7	
May 25, 1961	8.2	10.0	2.32 ± 0.11	5	
June 9, 1961	7.3	9.1	1.69 ± 0.06	5	
June 14, 1961	9.7	11.0	1.25 ± 0.06	20	
June 22, 1961	9.7	10.7	10.2 ± 0.2	20	
July 13, 1961	7.3	10.7	1.43 ± 0.09	20	
Aug. 3, 1961	7.3	11.0	0.36 ± 0.02	4	
Aug. 24, 1961	10.1	11.3	2.05 ± 0.06	4	
Sept. 1, 1961	9.8	11.7	0.89 ± 0.03	4	0
Sept. 5, 1961	8.8	10.7	1.37 ± 0.03	4	0
Sept. 7, 1961	6.7	7.9	0.77 ± 0.03	4	0
Sept. 11, 1961	9.8	11.3	1.01 ± 0.10	4	0.8
Sept. 15, 1961	13.1	13.4	0.91 ± 0.07	4	1.8
Sept. 15, 1961	11.9	13.4	0.67 ± 0.10	4	0
Sept. 18, 1961	13.4	13.7	3.5 ± 0.1	4	0.4
Sept. 19, 1961	11.9	13.4	1.55 ± 0.06	4	0.9
Sept. 22, 1961	12.8	14.0	2.24 ± 0.16	13	0
Sept. 22, 1961	9.1	14.0	1.19 ± 0.07	20	0
Sept. 25, 1961	8.8	10.4	0.66 ± 0.04	20	3.3
Oct. 2, 1961	7.9-8.5	10.1	0.60 ± 0.03	20	17
Oct. 5, 1961	9.4	10.7	1.01 ± 0.06	4	9.1
Nov. 16, 1961	10.4	12.6	1.51 ± 0.07	20	10
March 2, 1962	8.8	10.0	2.07 ± 0.07	4	16
May 1, 1962	8.5	10.5	1.73 ± 0.03	4	5.1
May 9, 1962	9.1	10.7	1.08 ± 0.04	4	12
May 11, 1962	7.6	9.6	1.38 ± 0.04	4	10
May 17, 1962	7.9	9.1	1.29 ± 0.03	4	18
May 24, 1962	7.0	8.7	1.48 ± 0.02	4	11

The method of sampling and analysis has been described by *Bleichrodt* [1962] and *Bleichrodt and van Abkoude* [1963a]. Only results of samples containing less than 20% of artificial <sup>7</sup>Be (see asterisked footnote) have been included. The branching ratio of <sup>7</sup>Be decay has been assumed to be 10.4% [*Poenitz and Devolpi*, 1973].

\*Artificial <sup>7</sup>Be in the samples taken after September 1, 1961, has been estimated according to *Bleichrodt and van Abkoude* [1963a].

†One standard deviation due to radioactivity counting.

factor of about 1.5 (the land area in these latitude bands ranges from 50 to 58%).

Some of the <sup>7</sup>Be deposited on the earth's surface might be brought down from the lower stratosphere by cloud systems penetrating the tropopause [*E. R. Reiter*, 1975]. This phenomenon leads to an overestimate of the deposition of tropospheric <sup>7</sup>Be and thus to an underestimate of the residence time.

In summary, the maximum possible error in the mean annual deposition results in an overestimation of the residence time by a factor of 2 for the latitude band 30°-80°N. For the other bands in Table 4 it is smaller.

#### MEAN TROPOSPHERIC RESIDENCE TIME

Table 4 shows minimum values for the mean tropospheric residence time for various latitude bands as obtained when curve I in Figure 1 is used to estimate the tropospheric content of <sup>7</sup>Be (Table 2). If the average tropopause height is assumed to be 9 km, which is certainly not too high [*Oort and Rasmusson*, 1971; *Crutcher*, 1969], a minimum for the tropospheric residence time of 22 days is obtained. Only for an average tropospheric concentration of <sup>7</sup>Be smaller than 1.5 times the average value observed at sea level (Figure 2) would the mean residence time have been smaller than the value of 7 days proposed by *Martell and Moore* [1974]. In the case of a maximum error in the estimated deposition of <sup>7</sup>Be (see discussion above) a minimal residence time of 12.5 days is obtained for

the latitude band 30°-80°N. For the other four latitude bands considered (Table 4) a maximum error in the deposition, estimated as explained above, yields a mean residence time of 13 days.

The value of  $6.6 \times 10^8$  atoms m<sup>-2</sup> (Table 2), calculated with curve II in Figure 1 for an average tropopause height of 11 km [*Oort and Rasmusson*, 1971; *Crutcher*, 1969], will approximate the average tropospheric content of <sup>7</sup>Be much better. The residence time thus appears to be 33 days for the latitude band 40°-60°N. With curve III in Figure 1 and a tropopause level of 11 km the tropospheric content is  $7.0 \times 10^8$  atoms m<sup>-2</sup> (Table 2), and the tropospheric residence time for the latitude band 40°-60°N is 35 days. This shows that uncertainties in the <sup>7</sup>Be

TABLE 2. Calculated Tropospheric Content of <sup>7</sup>Be

Assumed Tropopause Height, km	Tropospheric Content,* 10 <sup>8</sup> atoms m <sup>-2</sup>		
	Curve I	Curve II	Curve III
8	3.7	4.4	4.5
9	4.3	5.2	5.3
10	4.8	5.9	6.1
11	5.4	6.6	7.0
12	5.8	7.2	7.8

\*Curves in Figure 1 were used for calculations.

TABLE 3. Deposition of <sup>7</sup>Be at Various Locations in the Northern Hemisphere

Location	Geographic Coordinates	Sampling Period	Rainfall, <sup>a</sup> mm/yr	<sup>7</sup> Be Concentration in Rain, <sup>b</sup> 10 <sup>6</sup> atoms/l	Reference
Leningrad	60°N, 30°E	1964–1965 <sup>c</sup>	659	11.2	<i>Gedeonov and Rys'yev</i> [1969]
Milford Haven	52°N, 5°W	October 1959 through August 1961 <sup>d</sup>	1374	8.2	<i>Peirson</i> [1963]
Rijswijk	52°N, 4°E	November 1960 through August 1962 <sup>e</sup>	905	11.6	<i>Bleichrodt and van Abkoude</i> [1963b]
De Bilt	52°N, 5°E	1970–1975	723	6.0	<i>Ministry of Health and Environment</i> [1971–1976]
Chilton	51°N, 1°W	October 1959 through August 1961 <sup>d</sup>	838	7.2	<i>Peirson</i> [1963]
Lille	51°N, 3°E	1970, 1972	647	5.6–5.9 <sup>f</sup>	<i>Commission of the European Communities</i> [1972a, 1974]
Heidelberg	49°N, 9°E	June 1960 through May 1961	780	9.6	<i>Schumann and Stoepler</i> [1963]
Freiburg	48°N, 8°E	1965–1969	991	15.5	<i>Hartwig and Sittkus</i> [1972]
Rome	42°N, 13°E	1973–1974	697	9.6	<i>Commission of the European Communities</i> [1976]
Osaka	35°N, 136°E	1970–1973	1196	5.7	<i>Matsunami et al.</i> [1974]

<sup>a</sup>Average precipitation for sampling period.

<sup>b</sup>Ratio of total <sup>7</sup>Be deposition during sampling period of whole number of years and total precipitation.

<sup>c</sup>Dash means 1964 through 1965.

<sup>d</sup>Periods considered are October 1959 through September 1960 and September 1960 through August 1961.

<sup>e</sup>Periods considered are November 1960 through October 1961 and September 1961 through August 1962.

<sup>f</sup>In some monthly samples the activity of <sup>7</sup>Be was below a particular limit. The higher value is obtained by assuming the <sup>7</sup>Be concentration in these samples to correspond to the limit value, and the lower one by taking the <sup>7</sup>Be concentration to be zero.

TABLE 4. Mean Tropospheric Residence Time of <sup>7</sup>Be for Various Latitude Bands as Calculated With Curve I in Figure 1 for Different Tropopause Levels

Latitude Band	Average Annual Rainfall, mm	Average Annual Deposition, atoms/m <sup>2</sup> yr	Residence Time, days		
			9 km*	10 km*	11 km*
40°–60°N	800	7.2 × 10 <sup>9</sup>	22	24	27
40°–70°N	720	6.5 × 10 <sup>9</sup>	24	27	30
30°–60°N	790	7.1 × 10 <sup>9</sup>	22	25	28
30°–70°N	730	6.6 × 10 <sup>9</sup>	24	27	30
30°–80°N	690	6.2 × 10 <sup>9</sup>	25	28	32

\*Tropopause level.

concentration in the upper troposphere do not affect the final result appreciably. For the other latitude bands mentioned in Table 4 the calculated residence times are a little longer.

A mean tropospheric residence time of the order of 35 days is in agreement with previous estimates [cf. *Martell and Moore*, 1974]. It may be concluded that there is no experimental evidence to support the suggestion of *Martell and Moore* [1974] that the residence time of cosmic-ray-produced nuclides is as small as the value of less than 1 week, which was deduced for <sup>222</sup>Rn decay products.

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